

# Source parameters determined using borehole recordings

Kazutoshi Imanishi

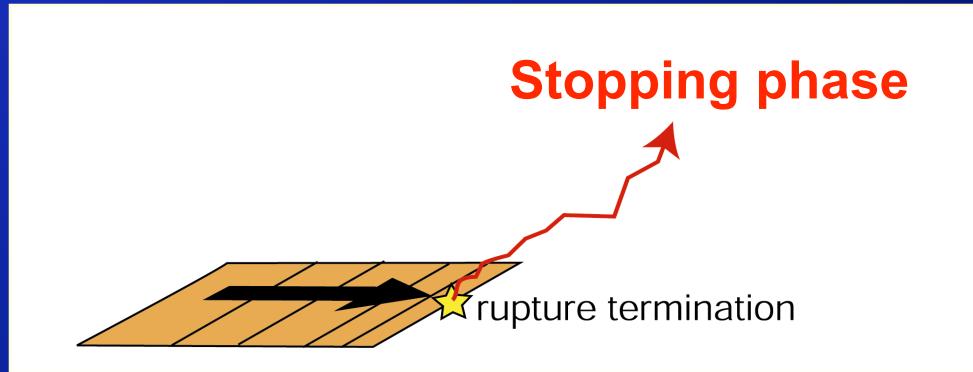
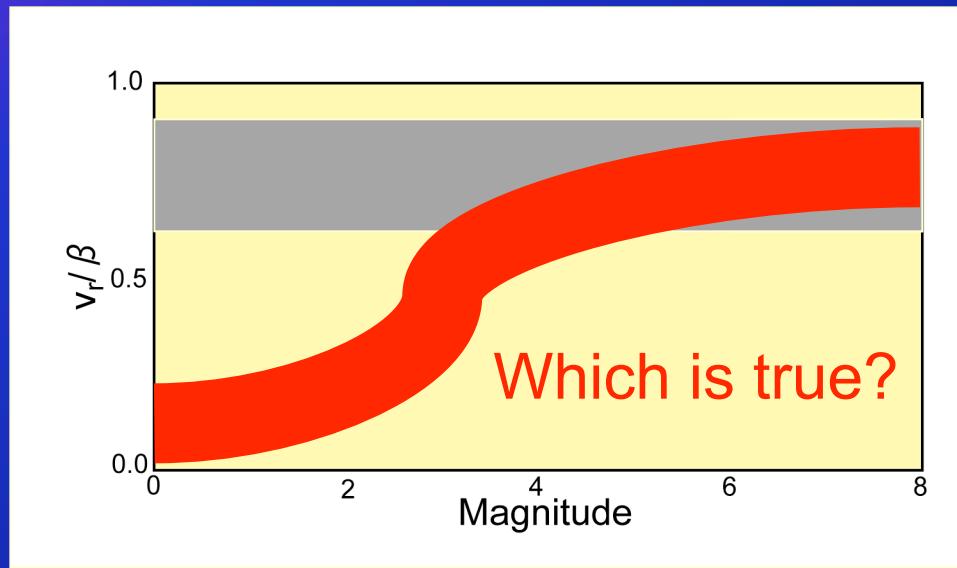
(Geological Survey of Japan, AIST)

1. Rupture velocities of microearthquakes determined using stopping phases
2. Earthquake source parameters determined by the SAFOD Pilot Hole vertical seismic array

# Rupture velocities of microearthquakes determined using stopping phases

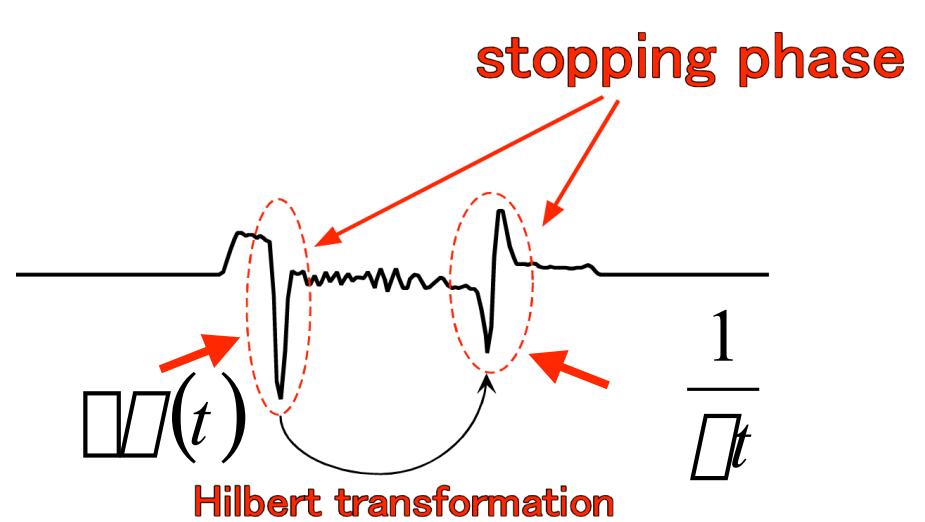
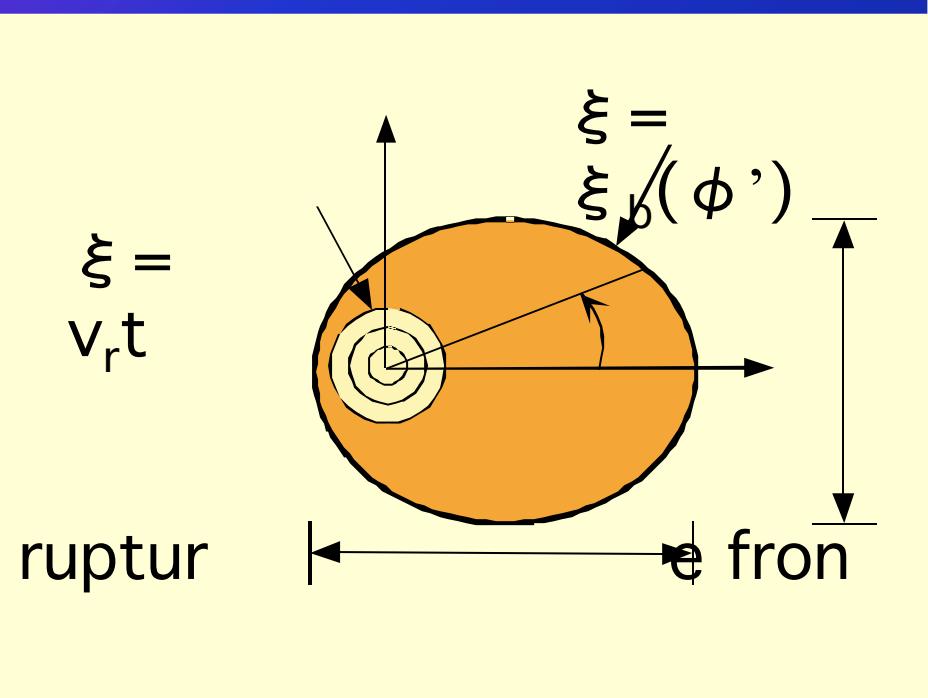
K. Imanishi<sup>1</sup>, M. Takeo<sup>2</sup>, H. Ito<sup>1</sup>, W. L. Ellsworth<sup>3</sup>,  
T. Matsuzawa<sup>2</sup>, Y. Kuwahara<sup>1</sup>, Y. Iio<sup>4</sup>, S. Horiuchi<sup>5</sup>,  
and Shiro Ohmi<sup>4</sup>

1. Geological Survey of Japan, AIST
2. ERI, University of Tokyo
3. U. S. Geological Survey
4. DPRI, Kyoto University
5. National Research Institute for Earth Science and Disaster Prevention

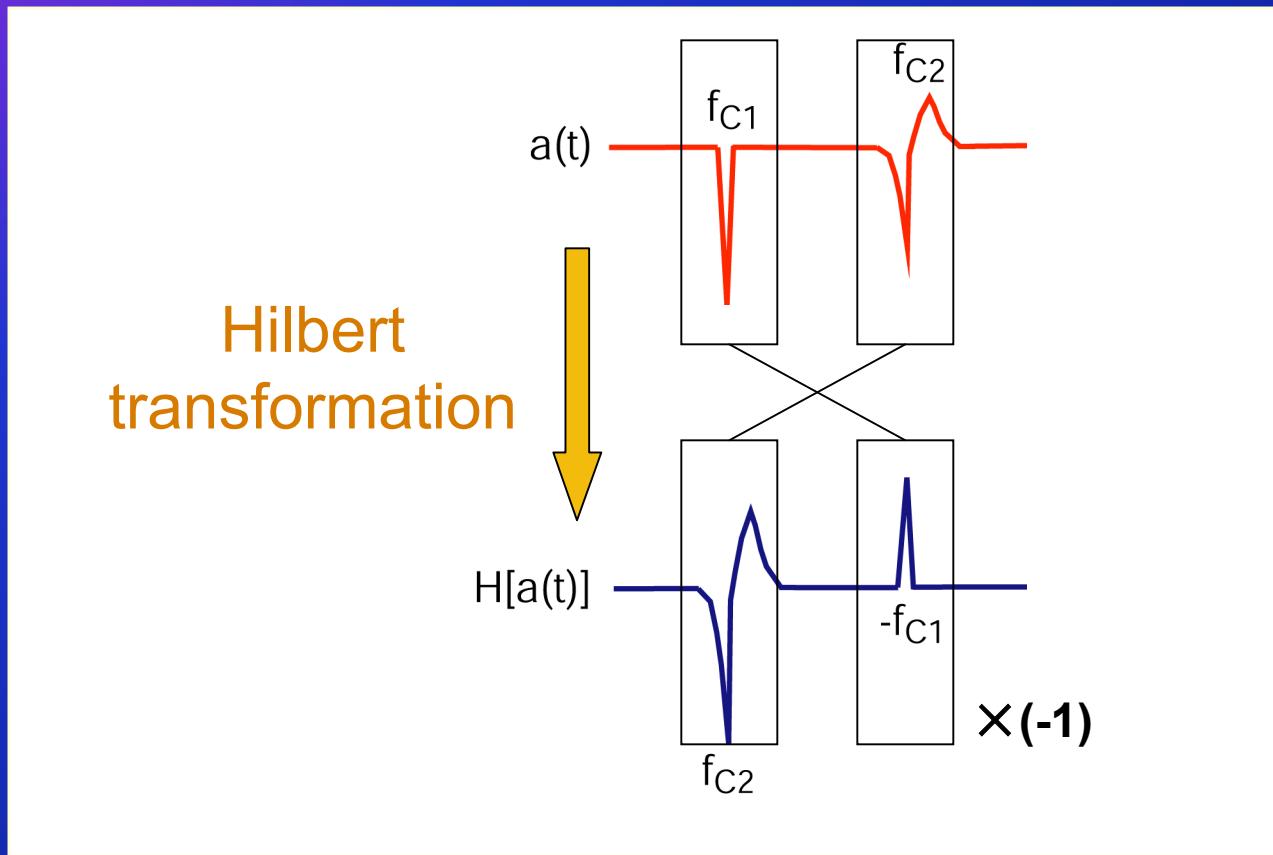


# What kind of stopping phases are radiated?

Elliptical fault model  
(Savage, 1966)



# Mutual relation of Hilbert transform pair



# Application to data

## Studied area

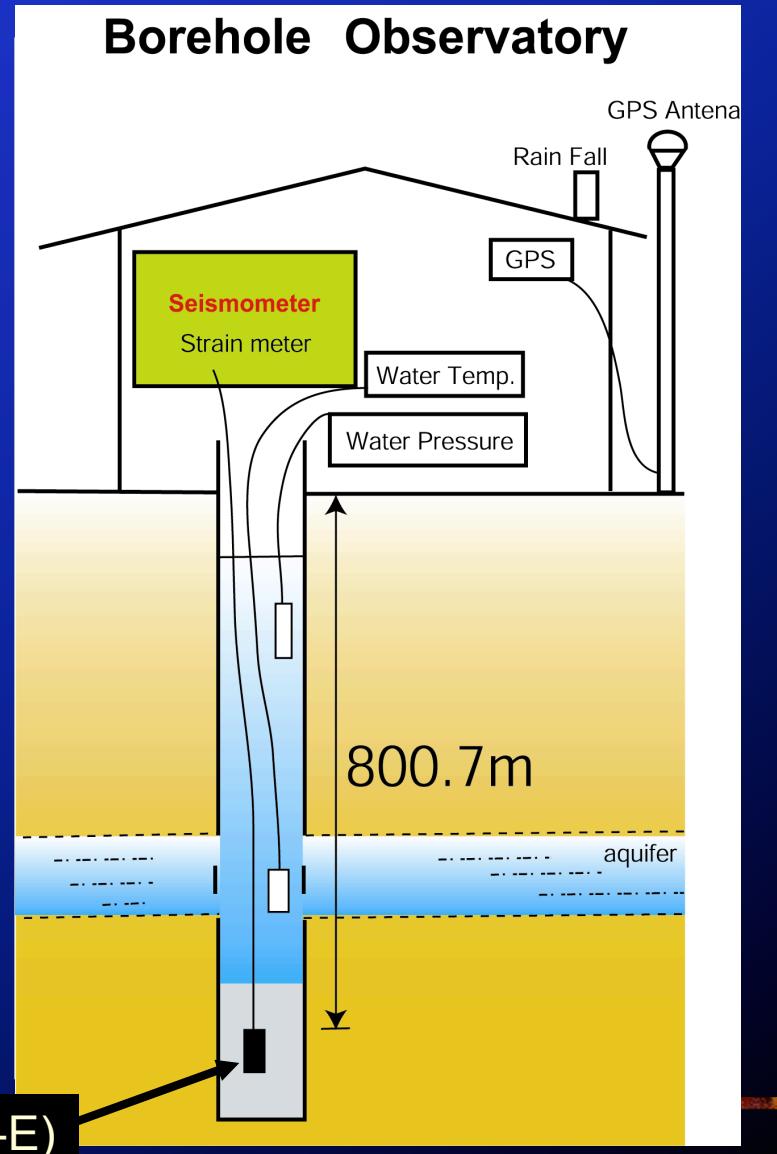
1. The area is characterized by persistent swarm activity.

(Most of earthquakes are less than M3.)

2. The Western Nagano earthquake (M=6.8) occurred in 1984.

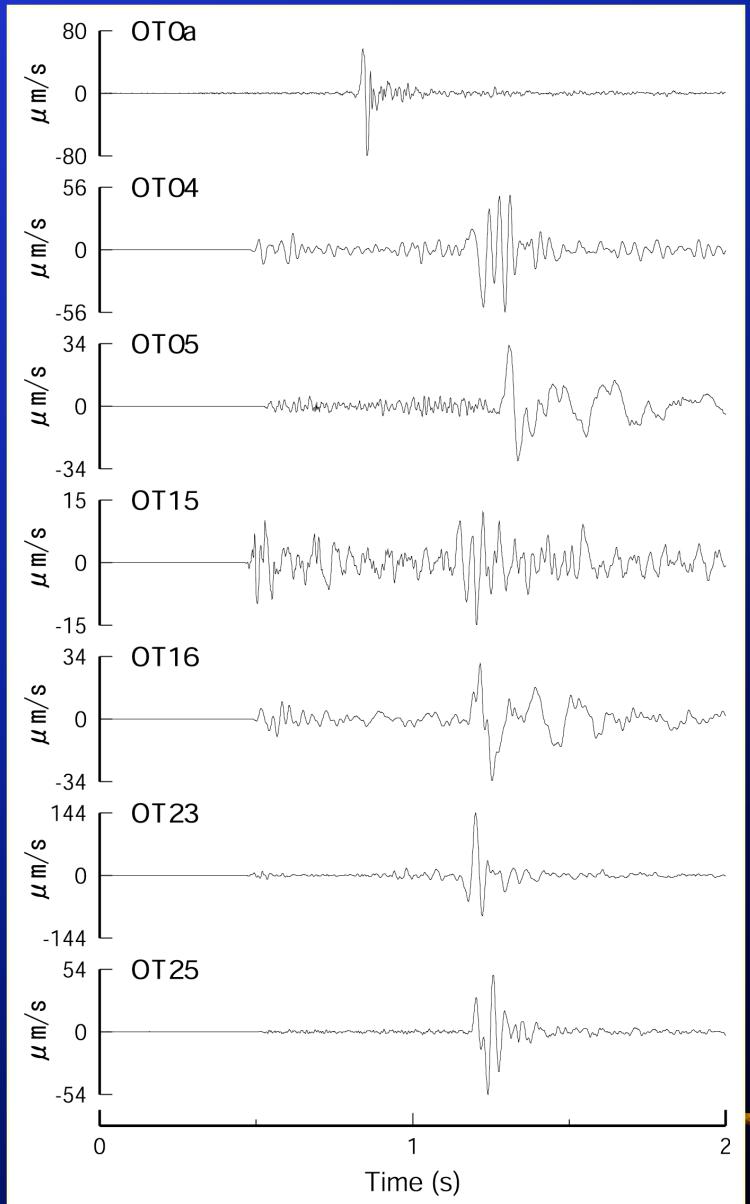
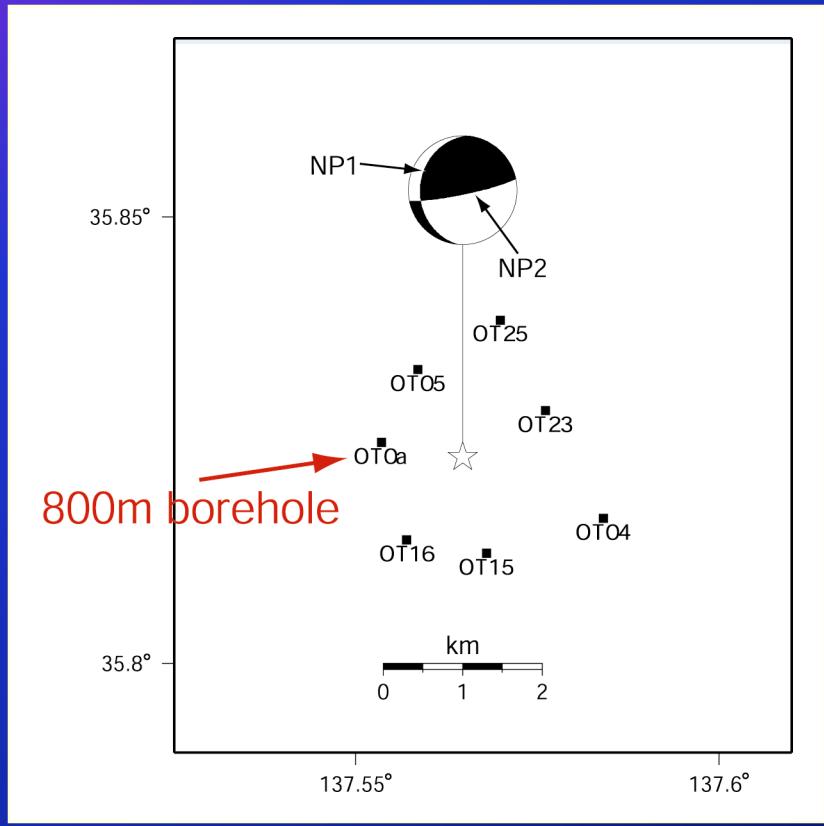
10 kHz sampling

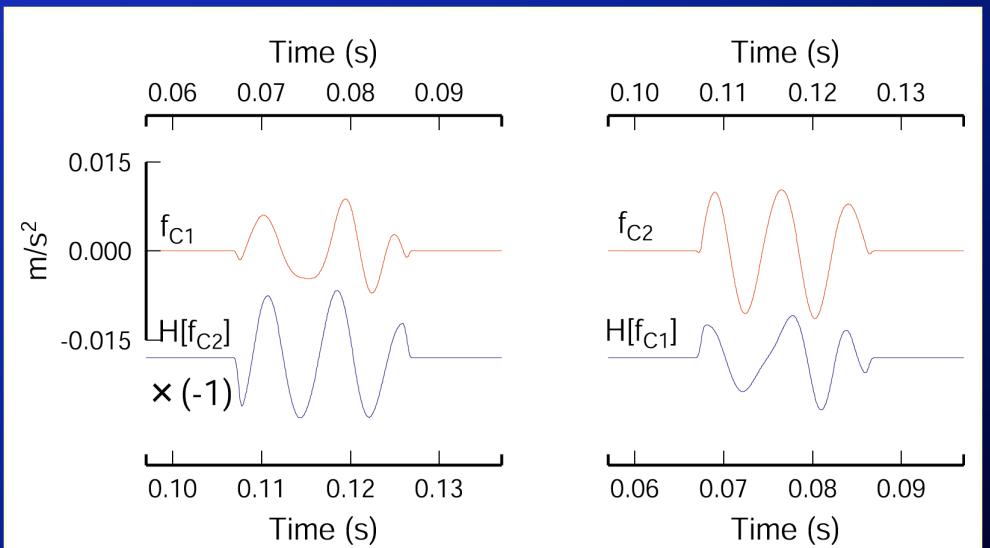
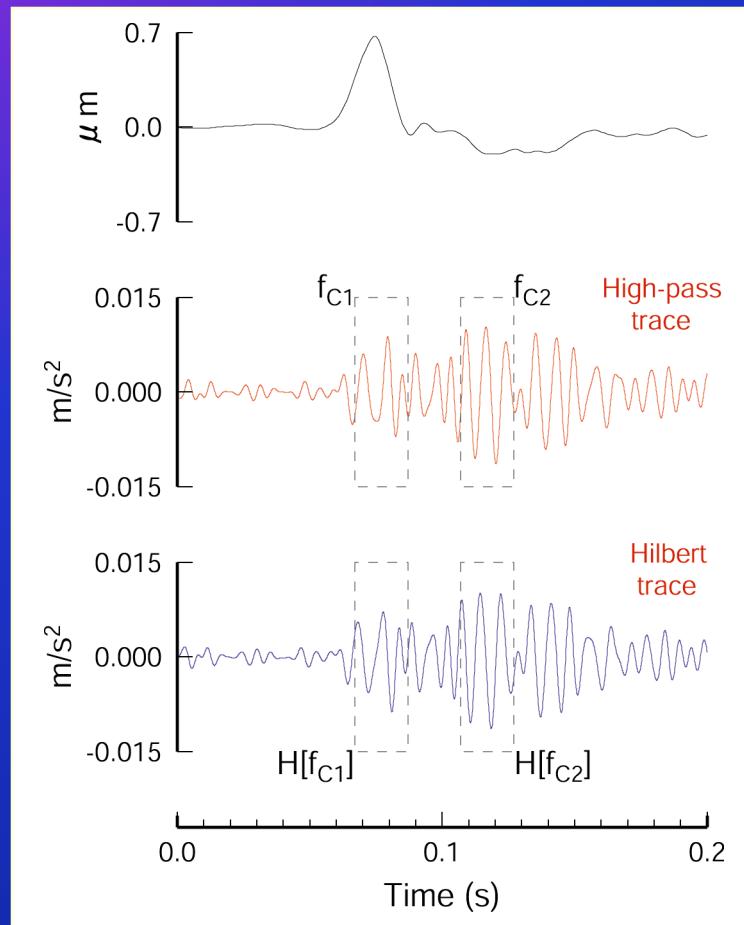
Three-component velocity transducer (L22-E)



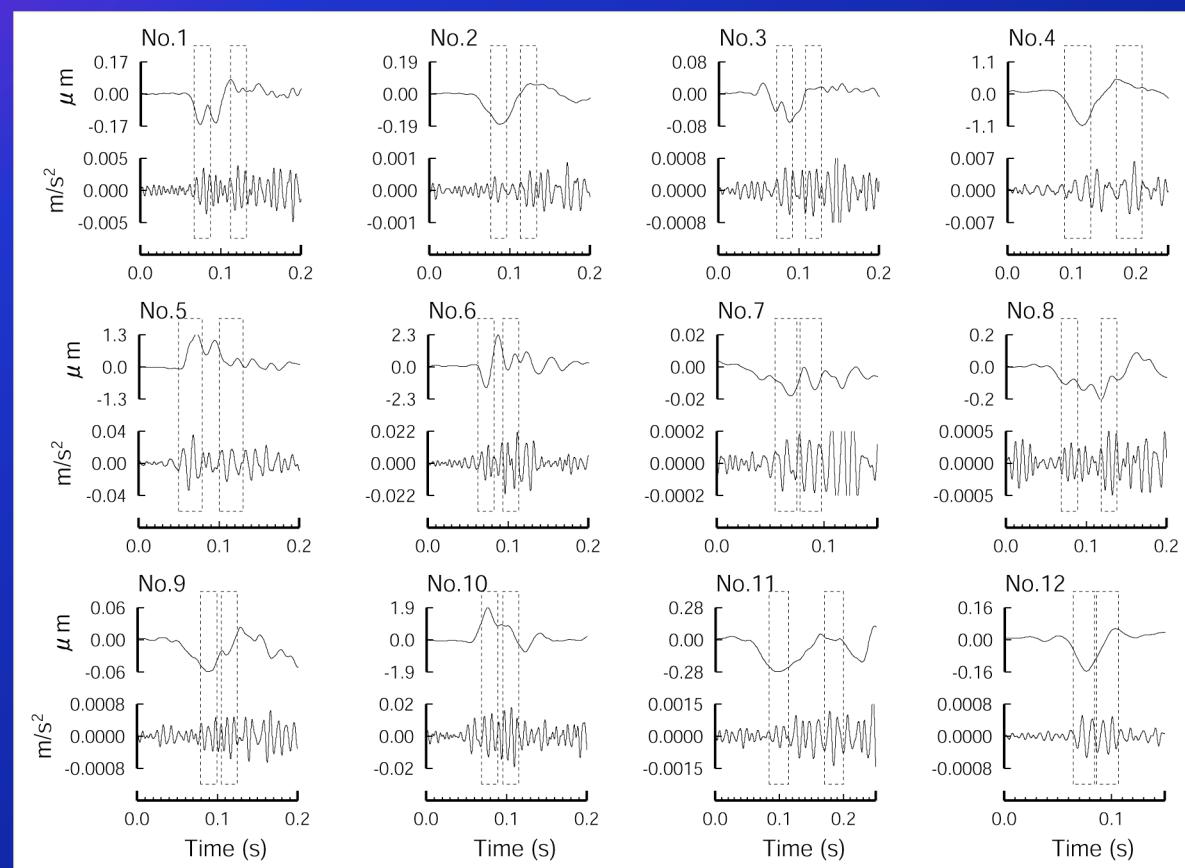
# Example

$M_w 1.4$





# Other earthquakes

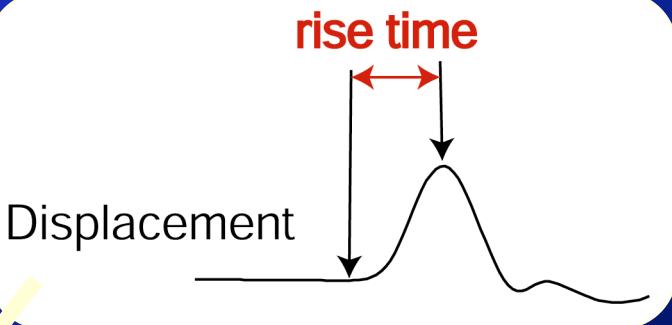


# Estimation of source parameters

800m borehole station

differential time of  
stopping phase

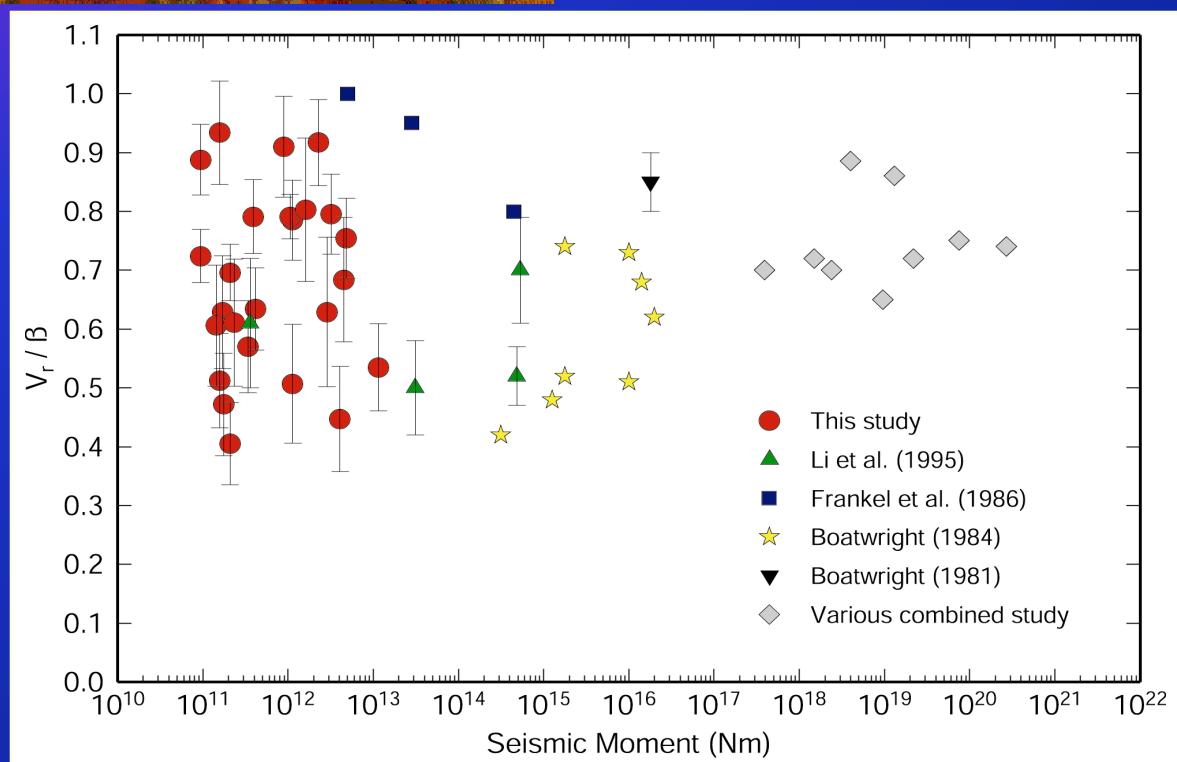
All stations



nonlinear inversion

rupture velocity, source size, rupture aspect ratio

# Rupture velocity vs seismic moment



Earthquakes are self-similar over a wide range of earthquake size and the dynamics of small and large earthquakes are similar.

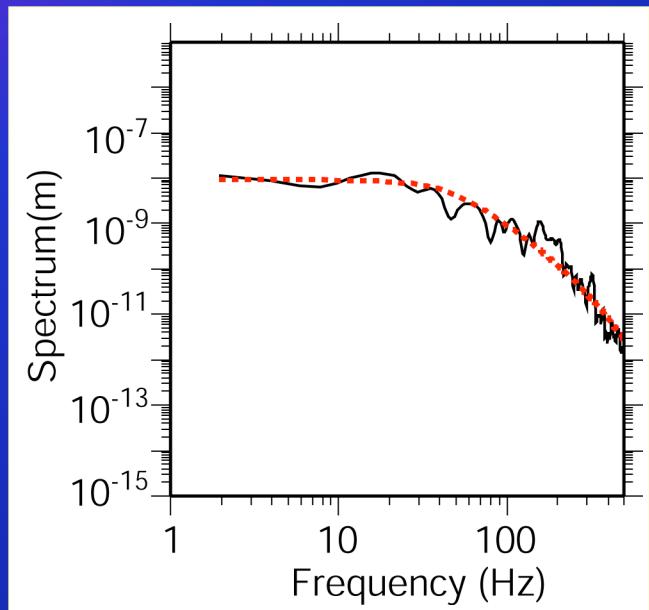
# Earthquake source parameters determined by the SAFOD Pilot Hole vertical seismic array

**K. Imanishi<sup>1</sup>, W. L. Ellsworth<sup>2</sup> and S. G. Prejean<sup>2</sup>**

1. Geological Survey of Japan, AIST

2. U. S. Geological Survey

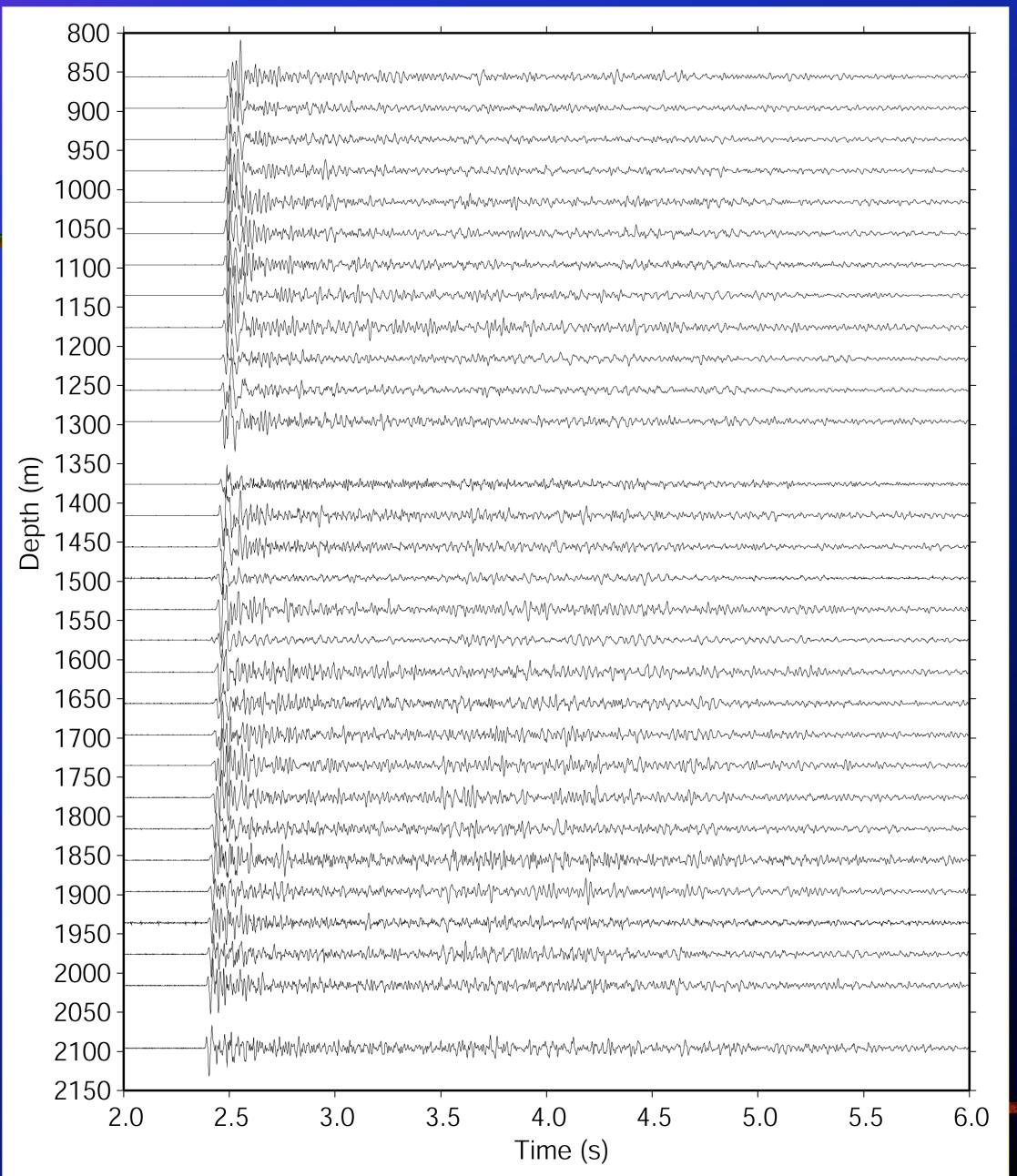
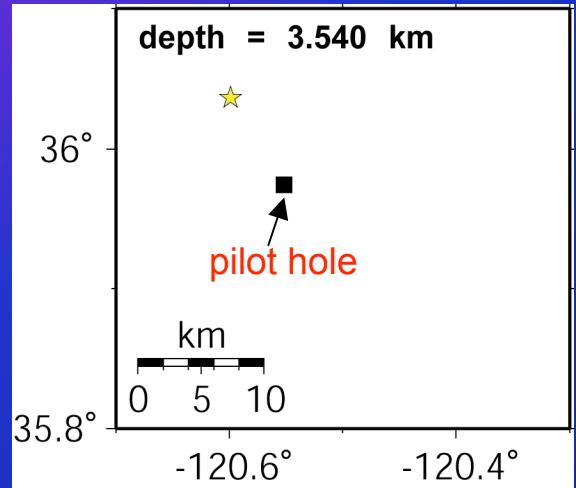
## spectrum analysis



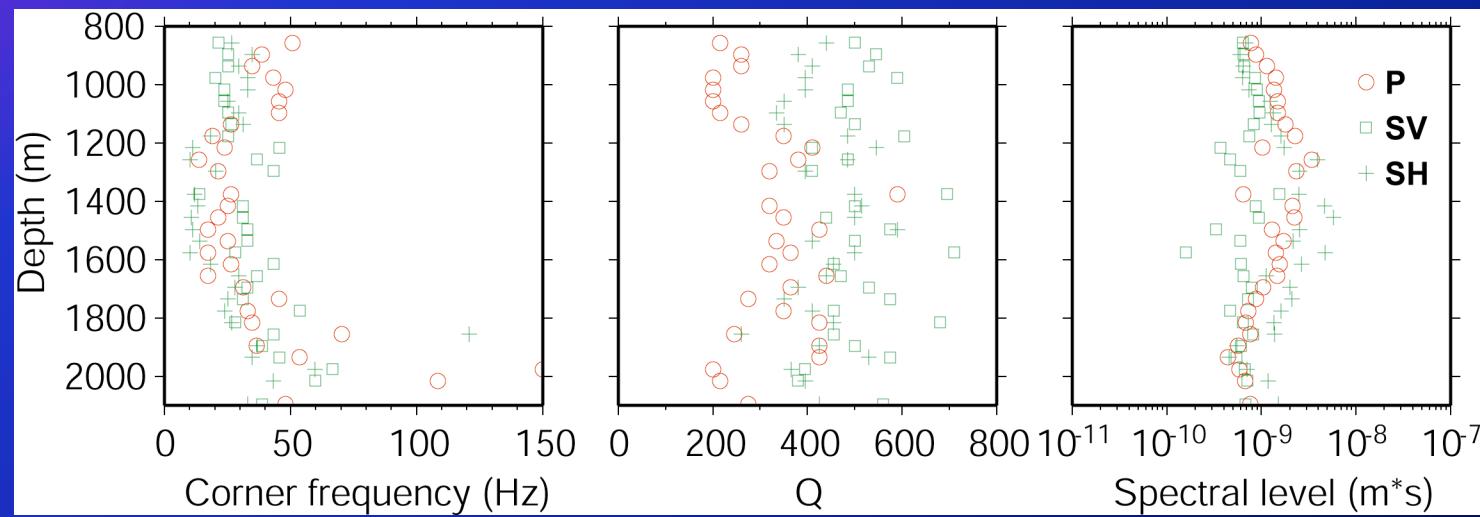
Corner frequency  
Q value  
Spectral level

Trade-off between  $f_c$  and Q.

## Example



## Source parameter estimations by fitting omega square model at each level.



There exists scatters in about only 1km difference of station locations.

## Smoothness constraint on $f_c$ , Q, and spectral level as a function of depth.

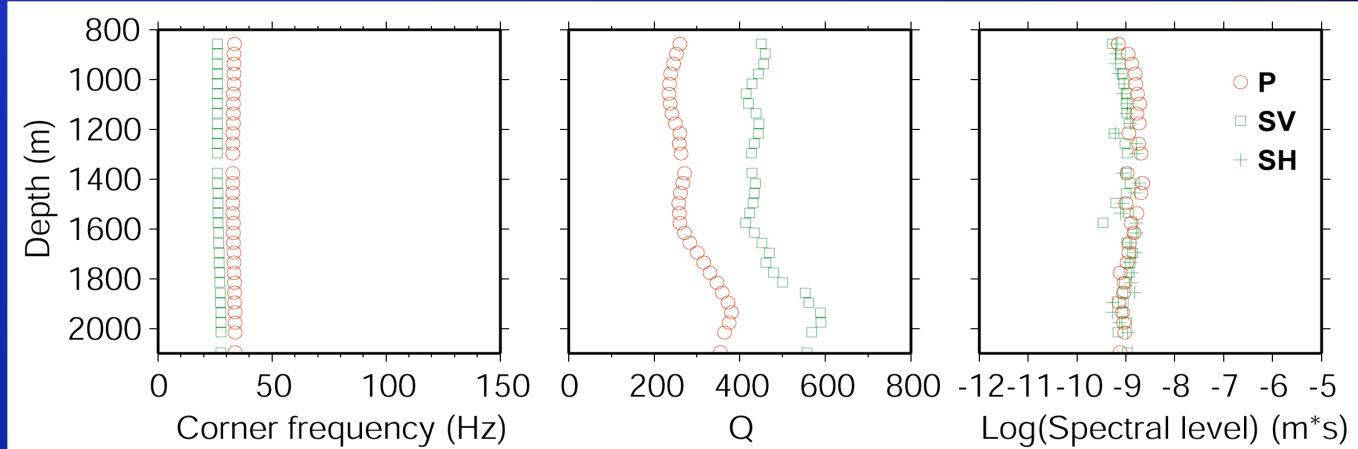
$$\left\| \frac{d \square g(m)}{\square_d} \right\|^2 + \square_1 \|D_1 m\|^2 + \square_2 \|D_2 m\|^2 + \square_3 \|D_3 m\|^2 = \min.$$

$d$  :data

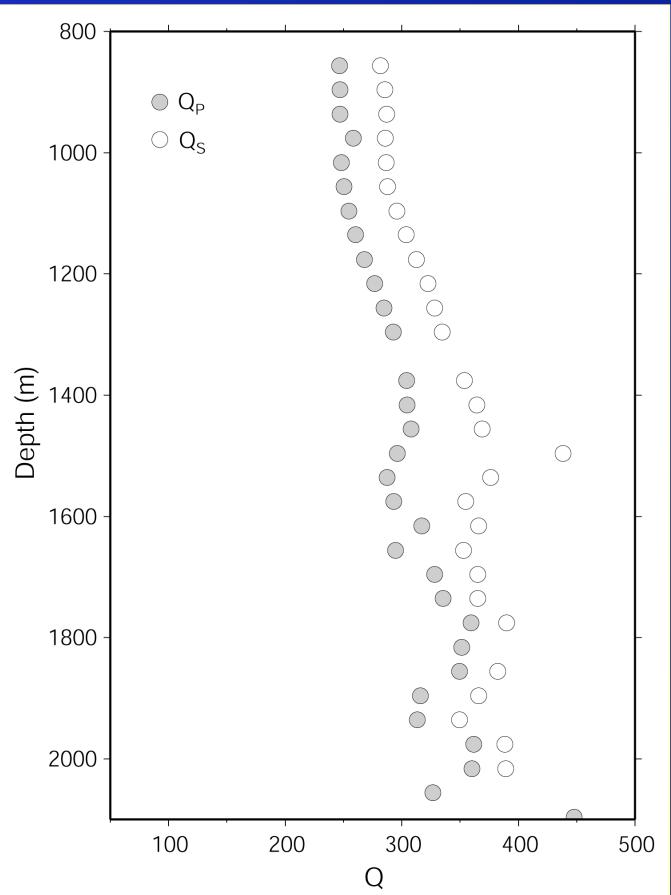
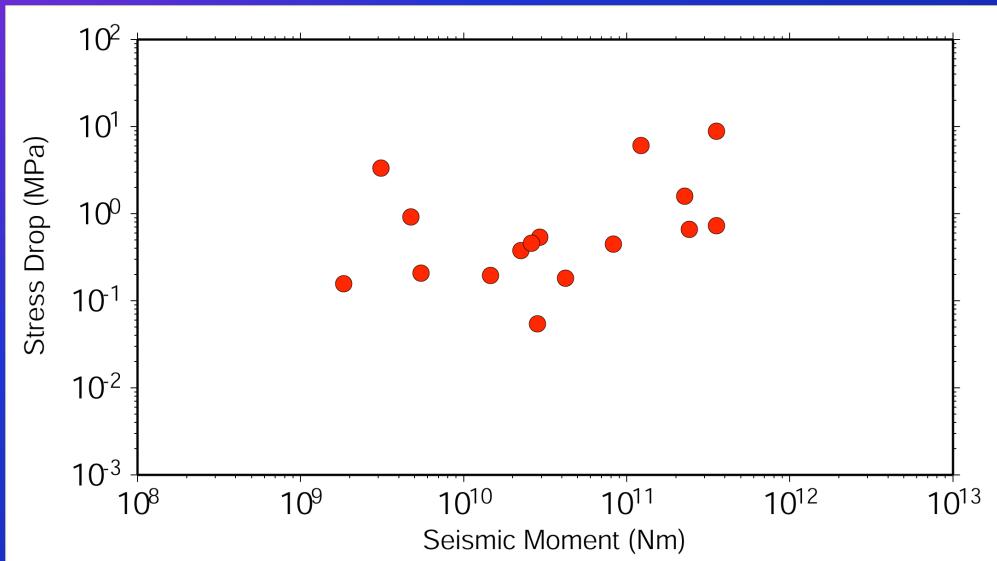
$g(m)$  :omega square model

$m$  :model parameter

$D$  :smoothing operator



# Result (Preliminary)



**Vertical borehole array is a powerful tool to determine source parameters of microearthquakes.**